







Cardiovascular Tissues

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(Nature, 2005)

Non-equilibration of hydrostatic pressure in blebbing cells

Guillaume T. Charras¹, Justin C. Yarrow¹, Mike A. Horton², L. Mahadevan^{1,3,4} & T. J. Mitchison¹





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Current models of the cytoplasm cannot account for spatiotemporal variations in hydrostatic pressure. We propose a new description of the cytoplasm based on poroelasticity. We consider cytoplasm to be composed of a porous, actively contractile, elastic network (cytoskeletal filaments, organelles, ribosomes), infiltrated with an interstitial fluid (...water, ions, soluble proteins), similar to a fluid-filled sponge. Contraction of the acto-myosin cortex creates a compressive stress on the cytoskeletal network, leading to localized increase in hydrostatic pressure & ... cytosol flow out of the network.....

Slow Stress Propagation in Adherent Cells

Biophysical J, 2008

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<u>ABSTRACT</u>: Mechanical cues influence...motility, differentiation, tumorigenesis.... study of how mechanical perturbations propagate across the cell is necessary to understand spatial coordination of cellular processes.

- Here we quantify magnitude & timing of *intracellular stress propagation*, using AFM and particle tracking by defocused fluorescence microscopy.
- The apical cell surface is locally perturbed by AFM cantilever indentation, and distal displacements are measured in 3 dimensions by tracking integrin-bound fluorescent particles.



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We observe an immediate response and slower equilibration, occurring over relaxation times that increase with distance from perturbation.

DISCUSSION

We compared our results to two material models: viscoelasticity and poroelasticity. A single-phase homogeneous viscoelastic material, such as the traditional spring-and-dashpot standard linear solid model, cannot explain the observed behavior, because it assumes that the material will simultaneously relax in response to a local perturbation with a single timeconstant. To determine if a heterogeneous viscoelastic model could explain this behavior, we modeled the experiment as a step-strain of a series of parallel spring-dashpot pairs (Voigt-Kelvin material (29,30)

 $T + \alpha \frac{dT}{dt} = E_1 e + \beta \frac{de}{dt} \cdot \cdot \cdot \cdot \cdot \cdot$

The poroelastic model can account for the observed slowdistance-dependent equilibration across the cell. The biphasic nature of a poroelastic material results in both a fast propagation of stress through the solid phase (cytoskeleton), and a much slower diffusive equilibration of hydrostatic pressure of the fluid phase (cytosol), resulting in increasing equilibration time with distance (20).





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 Our experimental results are not explained by traditional viscoelastic models of cell mechanics, but they are consistent with predictions from poroelastic models that include both cytoskeletal deformation and flow of the cytoplasm....

Musculoskeletal Tissues



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The role of viscoelasticity of collagen fibers in articular cartilage: axial tension versus compression



<u>Abstract:</u>For axial tension, collagen viscoelasticity was found to account for most of the stress relaxation, while the effects of fluid pressurization on the tensile stress were negligible. In contrast, for axial compression, the dominant mechanism for stress relaxation arose from fluid pressurization and fluid flow.....

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Kinetics of swelling of gels

Toyoichi Tanaka and David J. Fillmore

(J Chem Phys, 1979)

Gel displacement *u* satisfies a poroelastic diffusion equation!

$$\frac{\partial u}{\partial t} = D_{\text{gel}} \frac{\partial}{\partial r} \left[\frac{1}{r^2} \left(\frac{\partial}{\partial r} (r^2 u) \right) \right]$$
$$\tau_{\text{gel}} = \frac{L^2}{\pi^2 D_{\text{gel}}}$$

 $"D_{gel}" = "H" \bullet "k"$ H = (2G + λ) gel elasticity k = <u>gel hydraulic permeability</u>

(see Section 7.5 of text, page 260-261)



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J Applied Physics, 1954

JOURNAL OF APPLIED PHYSICS

VOLUME 26, NUMBER 2

FEBRUARY, 1955

Theory of Elasticity and Consolidation for a Porous Anisotropic Solid

Maurice Biot



The author's previous theory of elasticity and consolidation for isotropic materials [J. Appl. Phys. 12, 155-164 (1941)] is extended to the general case of anisotropy. The method of derivation is also different and more direct. The particular cases of transverse isotropy and complete isotropy are discussed.

1. INTRODUCTION

THE theory of consolidation deals with the settlement under loading of a porous deformable solid containing a viscous fluid. In a previous publication¹ a consolidation theory was developed for isotropic materials. The purpose of the present paper is to extend the theory to the most general case of anisotropy. The method by which the theory is derived is also more general and direct. The same physical assumption is introduced, that the skeleton is purely elastic and contains a compressible viscous fluid. The theory may therefore also be considered as a generalization of the theory of elasticity to porous materials. It is applicable to the prediction of the time bistory or stress and strain in a porous solid in which fluid scenage occurs. The general equations derived in Sec. 2 are applied to the case of transverse isotropy in Sec. 3. This is a case of particular interest in the application of the theory to soils and natural rock formations, since transverse isotropic is the type of symmetry usually acquired by rock under the influence of gravity. For an isotropic material the equations reduce to a simple form given in Sec. 4. They are shown to coincide with the equations derived in reference 1. Application of the theory to specific cases was made previously,2-4 and it was

sample of bulk volume V_b . It is understood that the term "porosity" refers as is customary to the effective porosity, namely, that encompassing only the intercommunicating void spaces as opposed to those pores which are sealed off. In the following, the word "pore" will refer to the effective pores while the sealed pores will be considered as part of the solid. It will be noted that a property of the porosity f is that it represents also a ratio of areas

$$j = S_p / S_b, \qquad (2.2)$$

i.e., the fraction S_p occupied by the pores in any crosssectional area S_b of the bulk material. It must beassumed, of course, that the pores are randomly disvibuted in location but not necessarily in direct the That this relation holds may be accertained by integrating S_p/S_b over a length of unity in a direction normal to the cross section S_b . The value of this integral then represents the fraction f of the volume occuried by the pores. It is seen that the ratio S_p/S_b is also independent of the direction of the cross section.

The stress tensor in the porous material is

$$\begin{cases} \sigma_{zz} + \sigma & \sigma_{zy} & \sigma_{zz} \\ \sigma_{yz} & \sigma_{yy} + \sigma & \sigma_{yz} \\ \sigma_{zz} & \sigma_{zy} & \sigma_{zz} + \sigma \end{cases}, \qquad (2.3)$$

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In the Jardin Darcy, Dijon, France

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$$\frac{\partial c_i}{\partial t} = D_i \frac{\partial^2 c_i}{\partial x^2}$$

Stress Relaxation 1110 = HL JU 80¹ U(x, t) B.C: 4 =0 @x=L incr. u,=u,@x,=0 I.c. u, = 0, t<0 00 $u_{1}\left(1-\frac{X_{1}}{L}\right)$ u (x, +) **Text** nsinh Eq. (7.55)described b 2 material prop. for all x, a

20.310J / 3.053J / 6.024J / 2.797J Molecular, Cellular, and Tissue Biomechanics $\ensuremath{\mathsf{Spring}}\xspace$ 2015

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