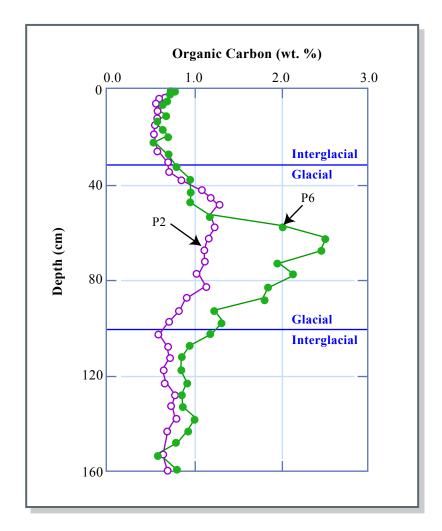
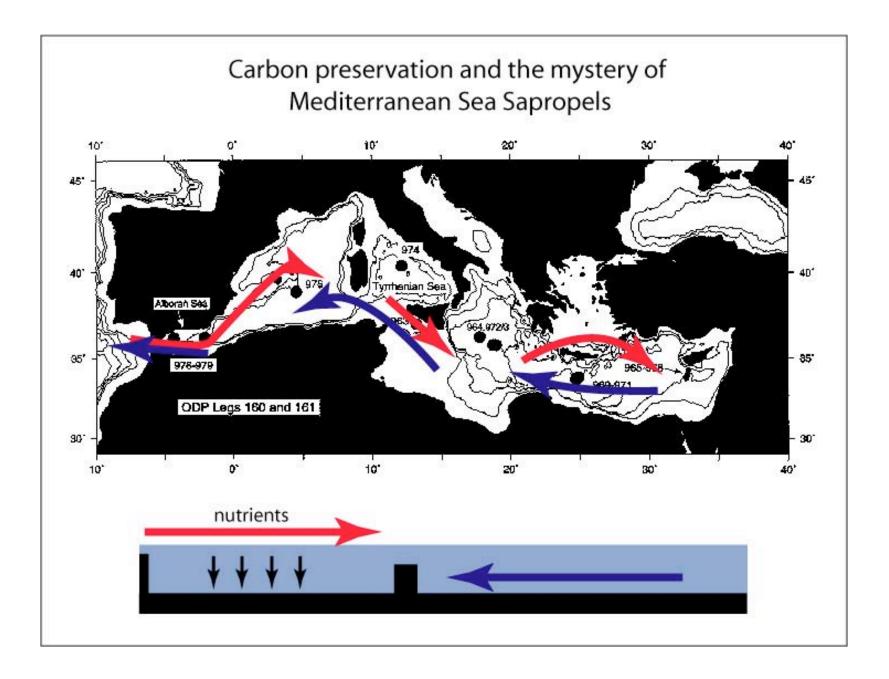
If the ocean is so efficient in remineralizing C and nutrients, why is C preserved at all in sediments?





Organic carbon with depth in late Pleistocene sediments from the eastern Equatorial Pacific Ocean. (Pederson and Calvert, AAPG Bull.(1990) v74, 454-466).



Sapropels

organic rich (2-14% TOC)

Periodic deposition

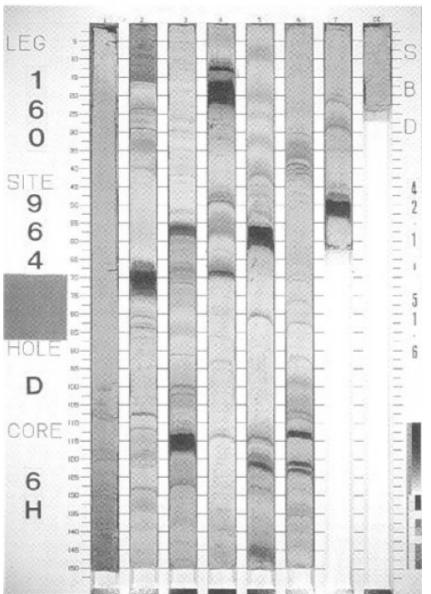
<1 cm to > 10 cm thick

Nonsapropels

Very organic lean (0.1% TOC)

Most of the deposition

Mediterranean Sea Sediment Core



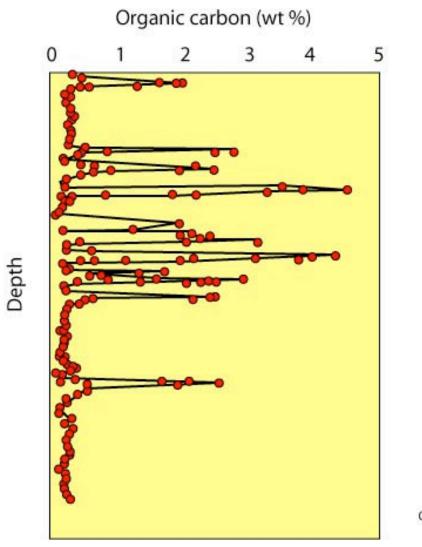
Sapropels

Any organic rich layer of sediment is called a sapropel, Sapropels in the Medditerranean Sea are very interesting however because the Med is one of the least productive bodies of water today, and sedimetns there are extremely depleted in organci carbon. A very long historical record of sapropel deposition was collected by the Ocean Drilling Program Legs 160 and 161 (see Initial reports...). Sapropels were first discovered in the Eastern Mediterranean Sea, but ODP found them to be synchronous in both basins. The shallowest sapropel is < 1m deep and can be sampled with a gravity core.

close-up of sapropel layer

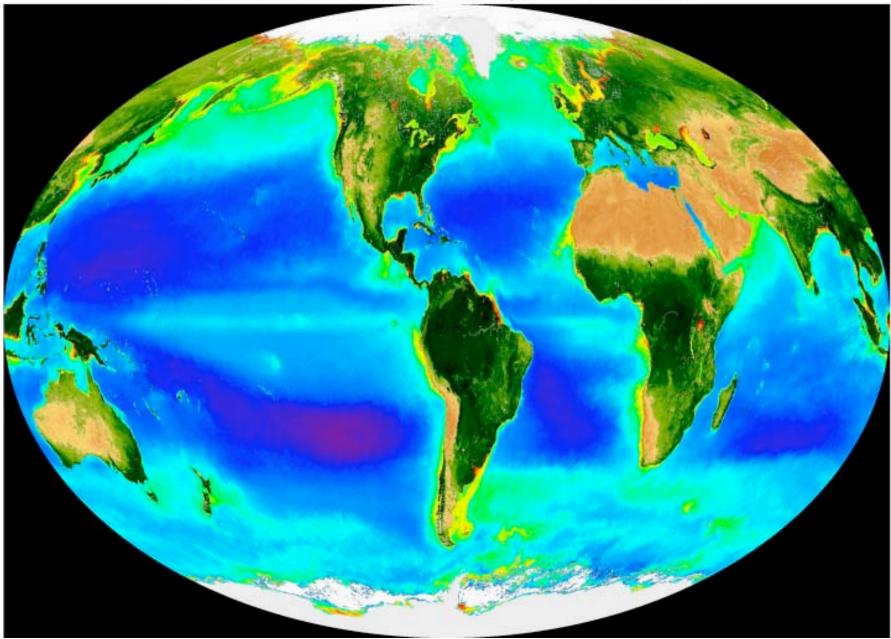


Organic carbon in Mediterranean Sea sediments

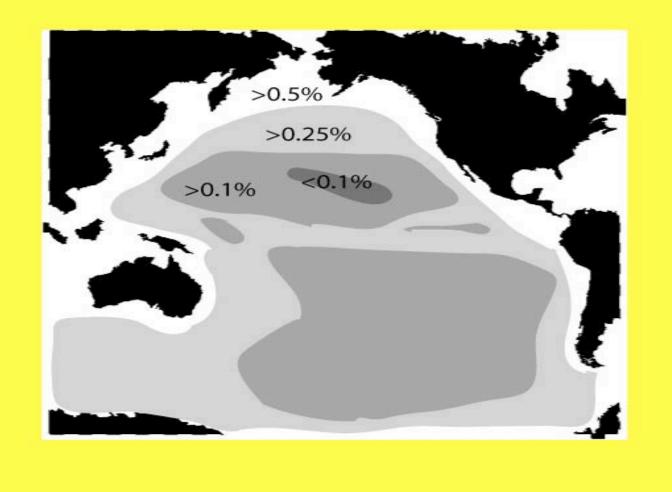


Calvert et al.

Global distribution of chlorophyll-a in September

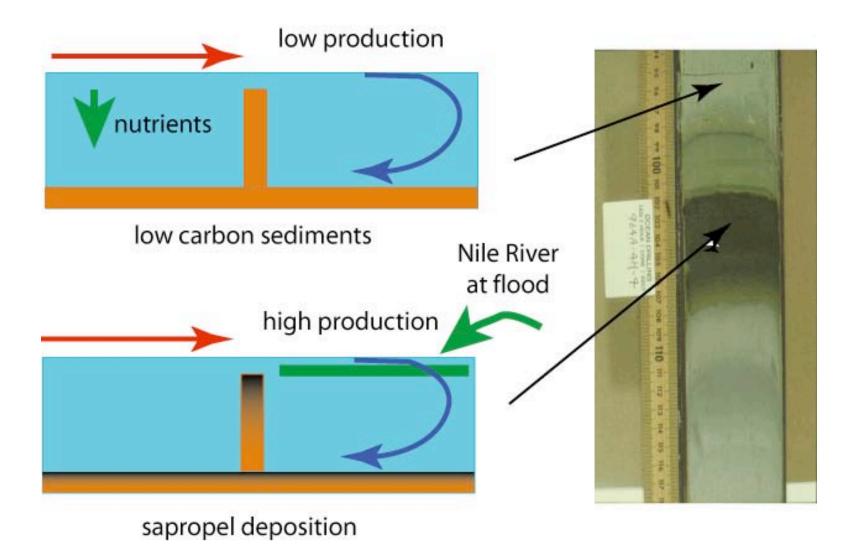


Organic carbon in surface sediments of the Pacific Ocean

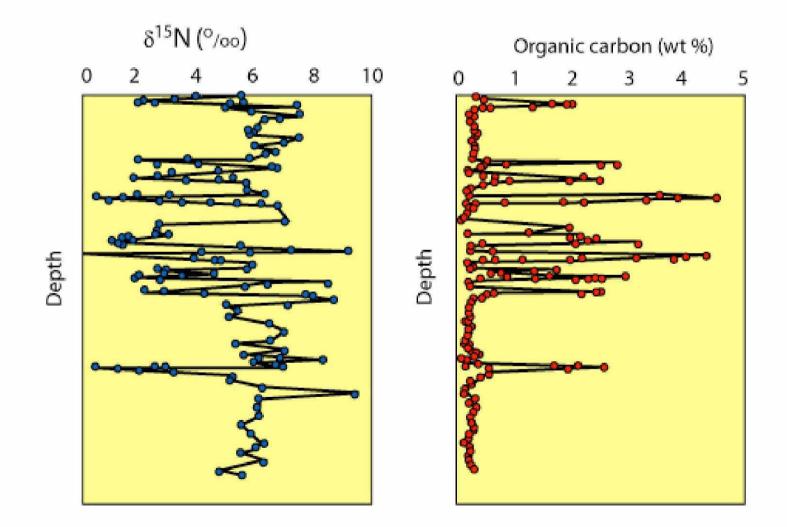


Calvert and Peterson 1998

Formation of Mediterranean Sea Sapropels Enhanced productivity hypothesis



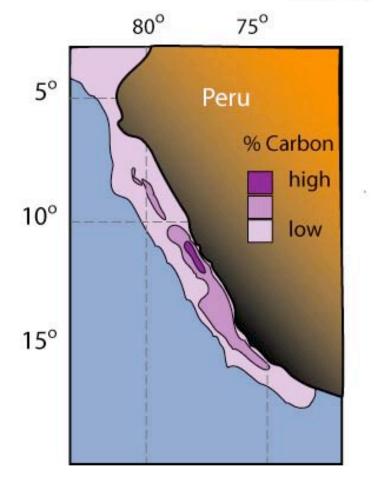
Correlation between %OC and nitrogen isotopes in Mediterranean Sea sediments

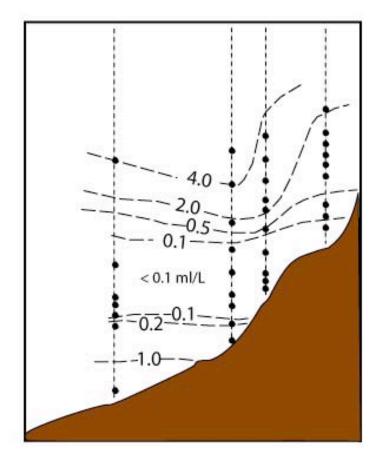


Nitrogen isotopes in paleoproductivity and denitrification studies

| Central Gyres | Upwelling Zones and Polar Seas | Oxygen Minimum Zones |
|--|---|--|
| Low Nutrients phytoplankton $\delta^{15}N = 5.9 \text{ per mil}$ $NO_3^- \longrightarrow N(\text{organic})$ $\epsilon = 0 \text{ per mil}$ | High Nutrients phytoplankton $\delta^{15}N < 5.9 \text{ per mil}$ $NO_3^- \longrightarrow N(\text{organic})$ $\epsilon = 6 \text{ per mil}$ | Low Oxygen phytoplankton $\delta^{15}N > 5.9 \text{ per mil}$ $NO_3^- \longrightarrow N_2$ $\epsilon = 20 \text{ per mil}$ |
| | Deep Sea Nitrate δ^{15} N = 5.9 per mil | |

Carbon accumulation and the oxygen minimum zone of the Peru upwelling system



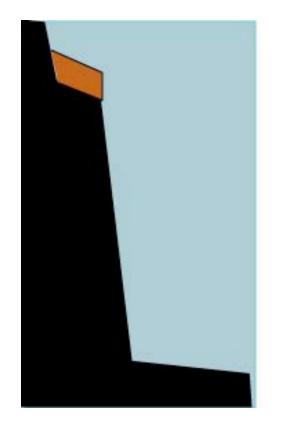


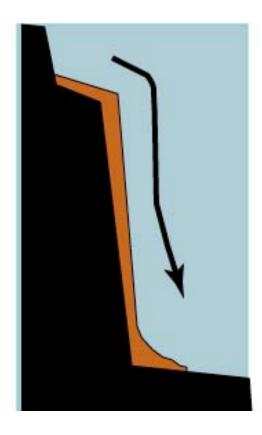
The effect of oxygen on carbon preservation in Maderia Abyssal Plain Turbidites

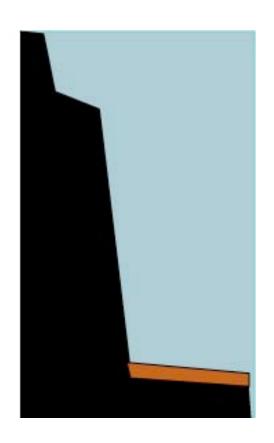
Before...

During...

and Voila!







The effect of oxygen on carbon preservation in Maderia Abyssal Plain Turbidites

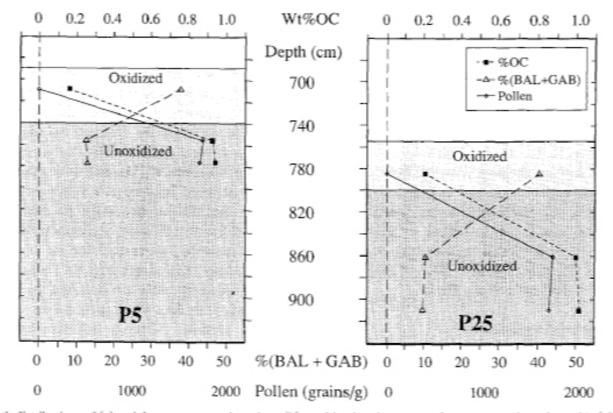
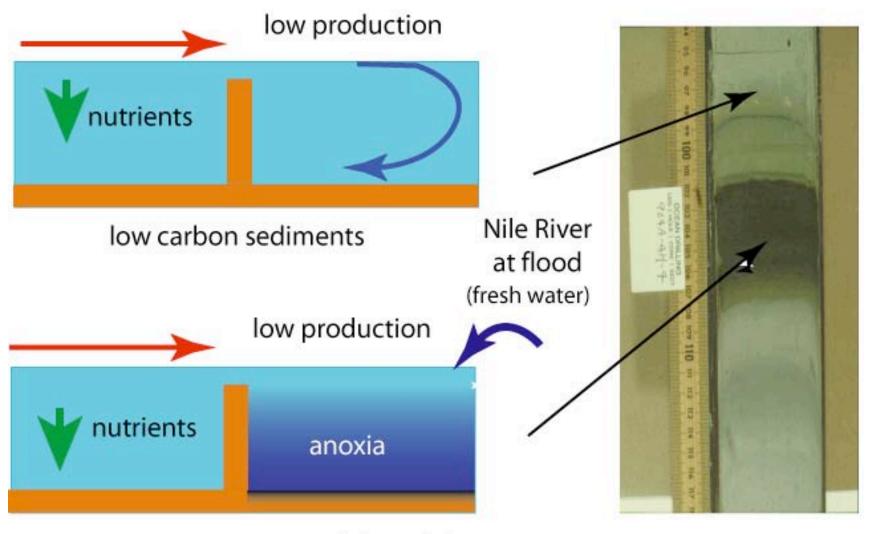


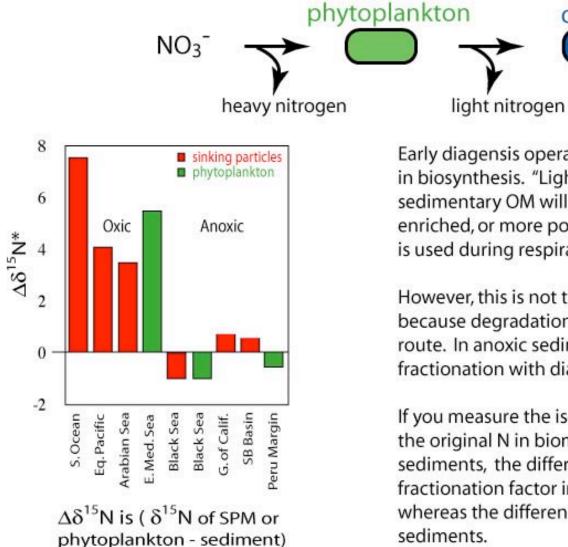
Fig. 10. Depth distributions of (a) weight percent organic carbon, (b) combined mole percent of two nonprotein amino acids (β -alanine plus γ -aminobutyric acid), and (c) total pollen abundances (grains g^{-1}) in oxidized and unoxidized sediments from two cores of the f-turbidite collected at separate sites in the Madeira Abyssal Plain (data from Cowie et al., 1995; Keil et al., 1994b).

Formation of Mediterranean Sea Sapropels - anoxia hypothesis



sapropel deposition

N isotope fractionation and early diagenesis



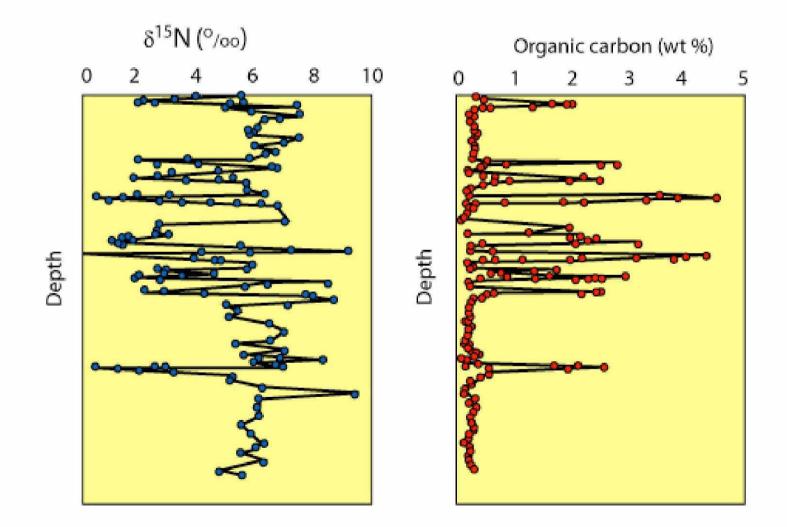
Early diagensis operates like N uptake in biosynthesis. "Light" N is used first, so that sedimentary OM will get heavier (more enriched, or more positive) in N-15 as nitrogen is used during respiration.

detritus

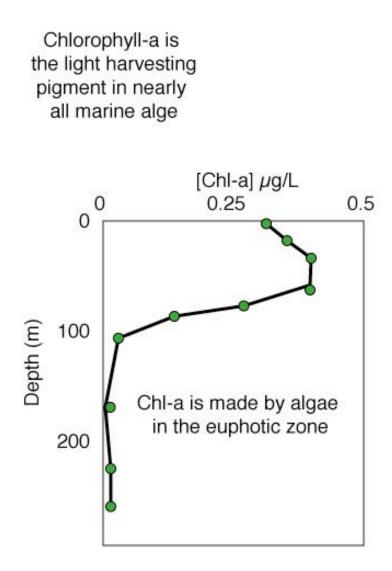
However, this is not the case in anoxic systems because degradation proceeds via a different route. In anoxic sediments, there is little N fractionation with diagenesis.

If you measure the isotopic difference between the original N in biomass, and N that is left in sediments, the difference will be large (large fractionation factor in N uptake) in oxic sediments, whereas the difference will be small for anoxic sediments.

Correlation between %OC and nitrogen isotopes in Mediterranean Sea sediments



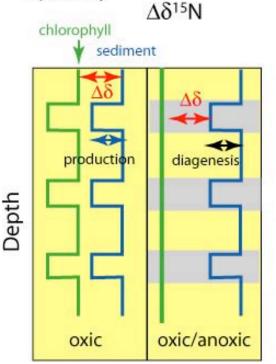
δ^{15} N in chlorophyll and sediment from Mediterranean Sea sapropels

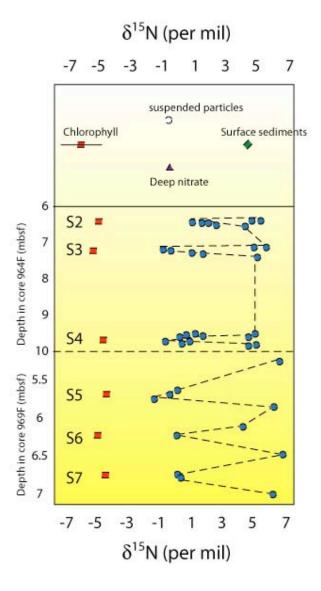


 Chlorophyll (Chl) will record the N isotope value of algal organic matter at the time of synthesis. The isotopic difference beteen Chl and algal biomass is a constant (5.9 per mil).

Dlagenesis acts on total N, but will not change the isotopic value of Chl.

4. If the isotopic difference ($\Delta\Delta$ 15N) decreases during sapropel formation, then N isotopic value is changing due to diagensis. If it is constant, then the value of N in the water column is changing and Δ N15 is set by changes in productivity





What causes sapropels to form in the Mediterranean Sea? And more generally, what processes act to preserve carbon In marine sediments?

close-up of sapropel layer



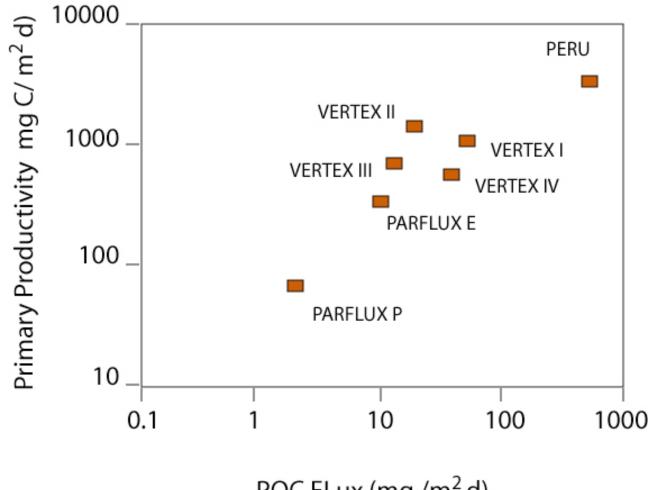
Enhanced productivity due to Inputs of nutrients?

or

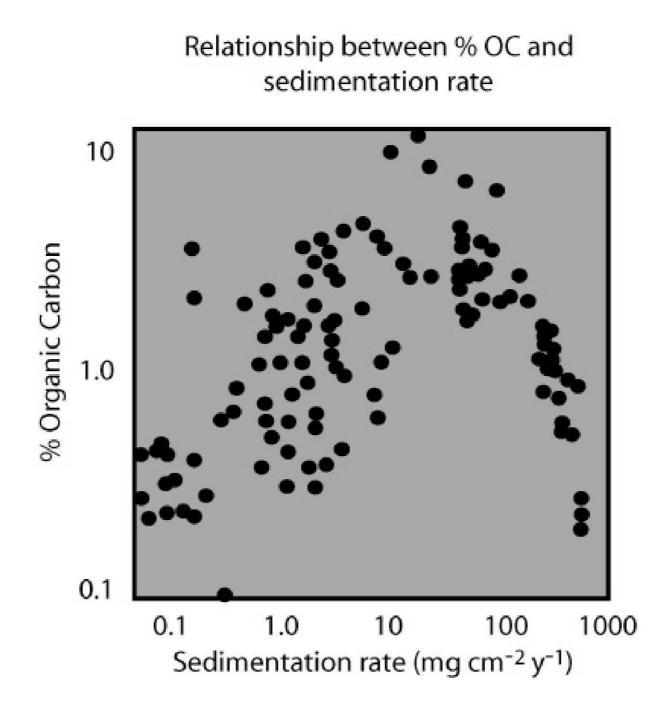
Enhanced preservation due to Bottom water anoxia?

The productivity vs preservation debate

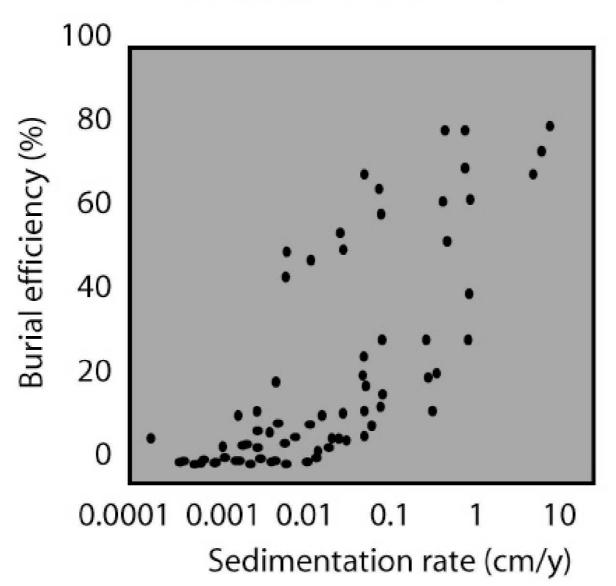
CP factor #1. The argument for productivity.....



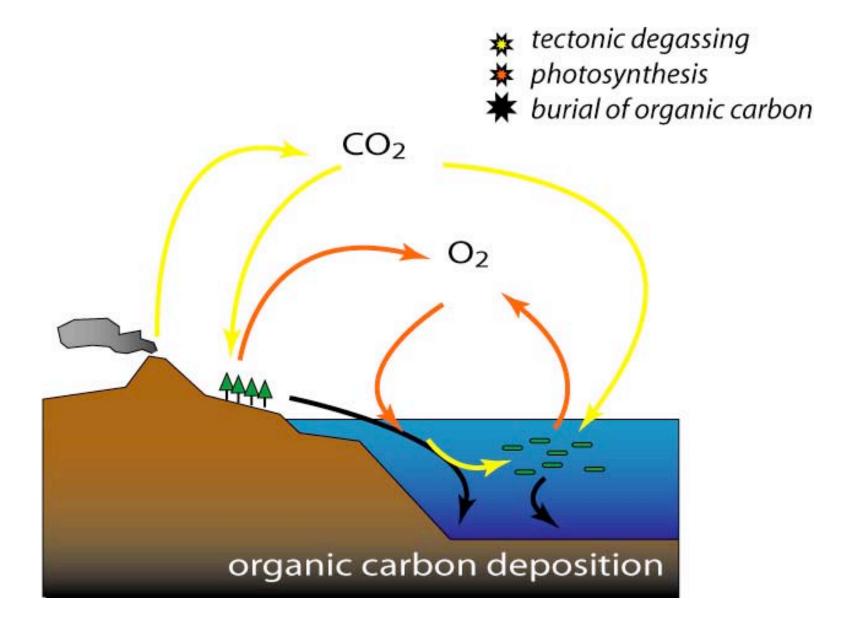
POC FLux (mg /m² d)



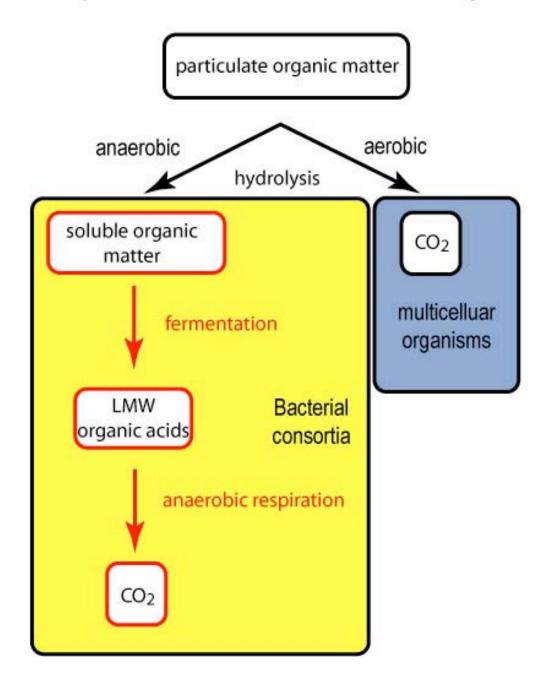
Relationship between burial efficiency and sedimentation rate



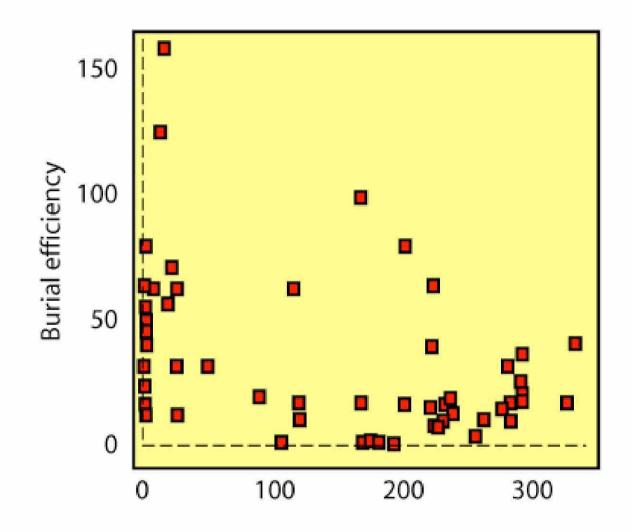
Coupling and feedbacks between carbon and oxygen cycles



Comparison of aerobic and anaerobic degradation

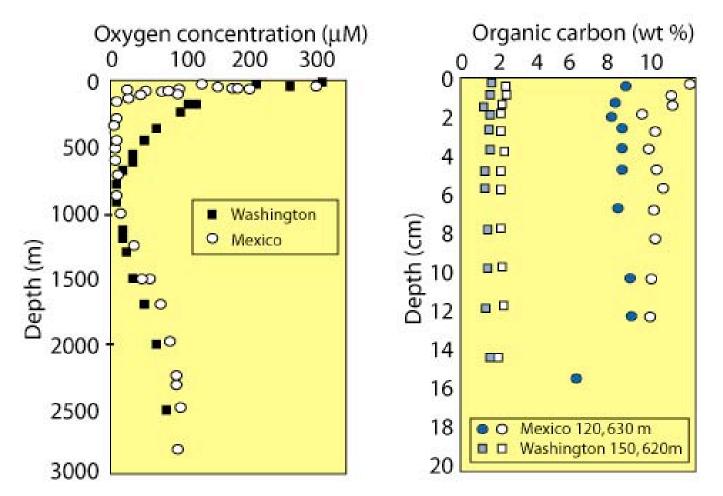


Effect of bottom water oxygen on burial efficiency



Bottom water oxygen (M)

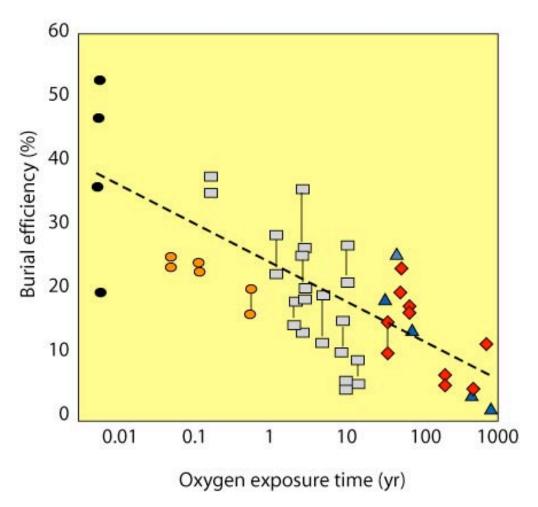
The effect of oxygen on carbon preservation in continental margin sediments



Hartnett et al. (1998) Nature v391, 572-574

The effect of oxygen has been refined somewhat to adjust for differences in exposure time, which is related to sedimentation rate (depth of O_2 penetration/sedimentation rate) = OET

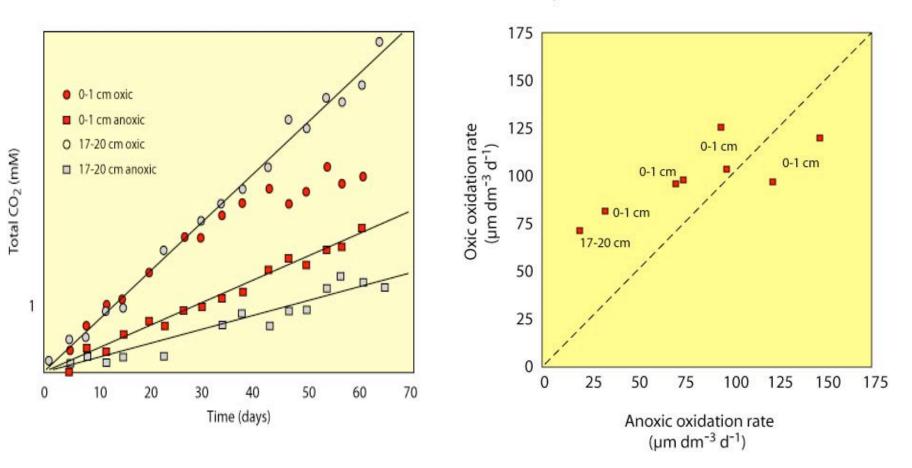
Effect of oxygen exposure time on burial efficiency



Is carbon more efficiently respired under oxic or anoxic conditions ?

Respiration of carbon in 0-1 cm and 17-20 cm sediment under oxic and anoxic conditions

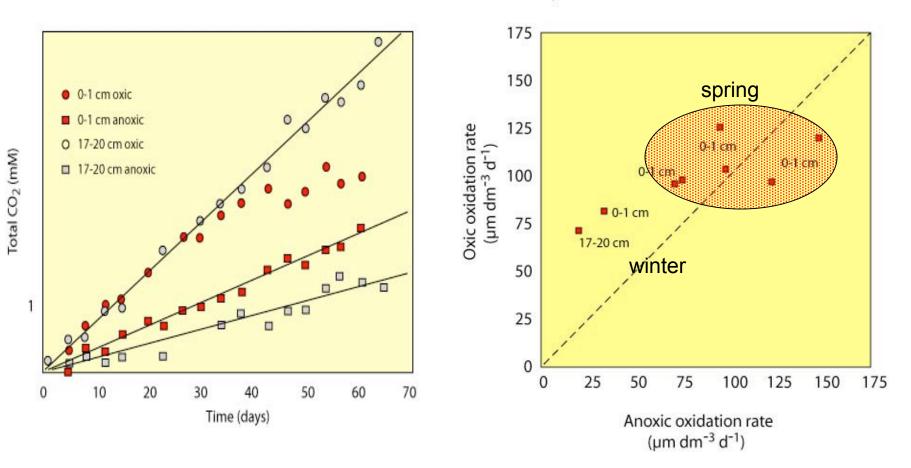
Comparison of oxic and anoxic degradation rates in surface and deep sediments



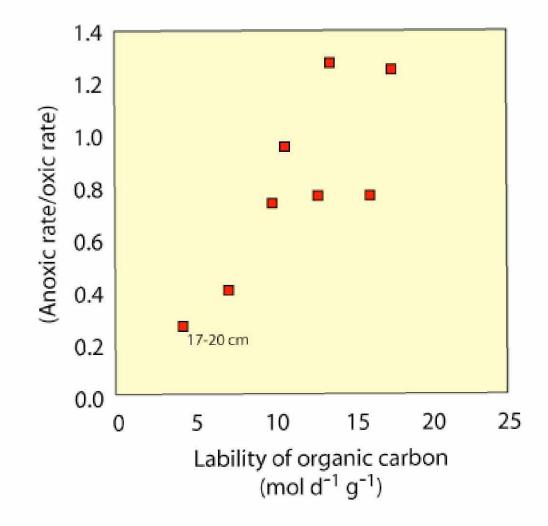
Is carbon more efficiently respired under oxic or anoxic conditions ?

Respiration of carbon in 0-1 cm and 17-20 cm sediment under oxic and anoxic conditions

Comparison of oxic and anoxic degradation rates in surface and deep sediments

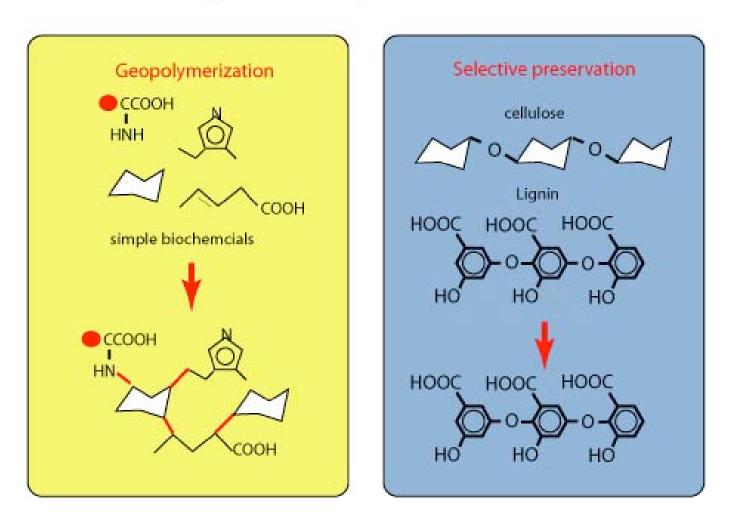


Oxidation rate and the lability of organic carbon

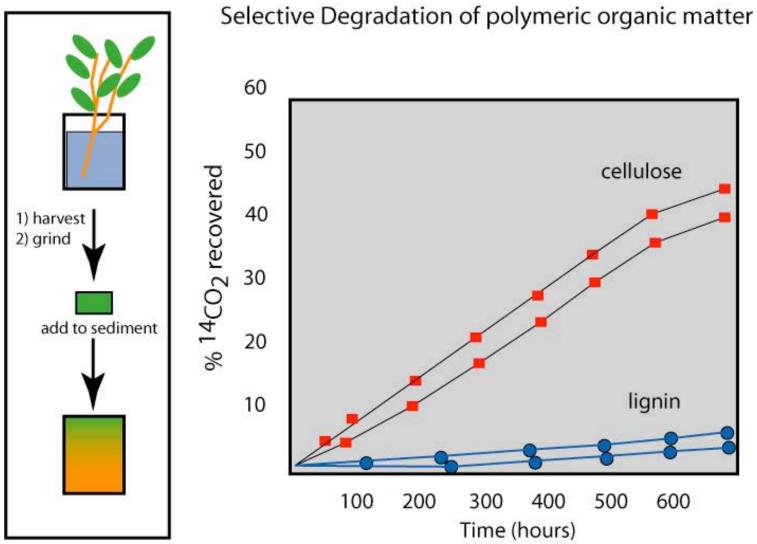


CP Factor #3 The composition of organic matter

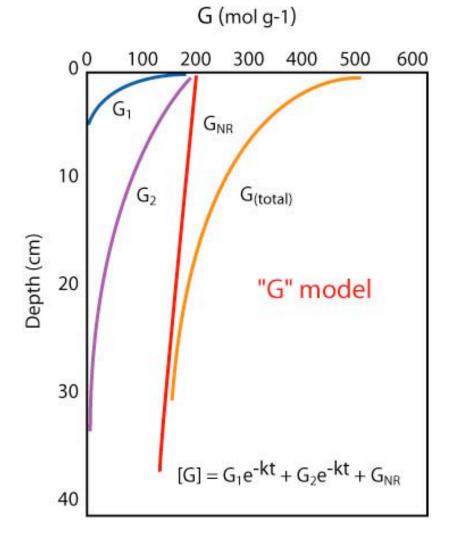
Effect of chemical composition on organic matter degradation in sediments

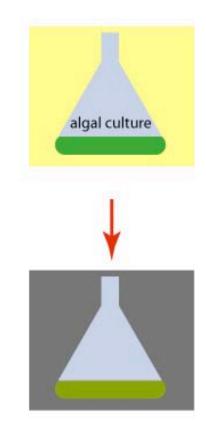


Was the composition of OM in anoxic sediments different than the OM in oxic sediments, thereby producing the **Observed difference?**

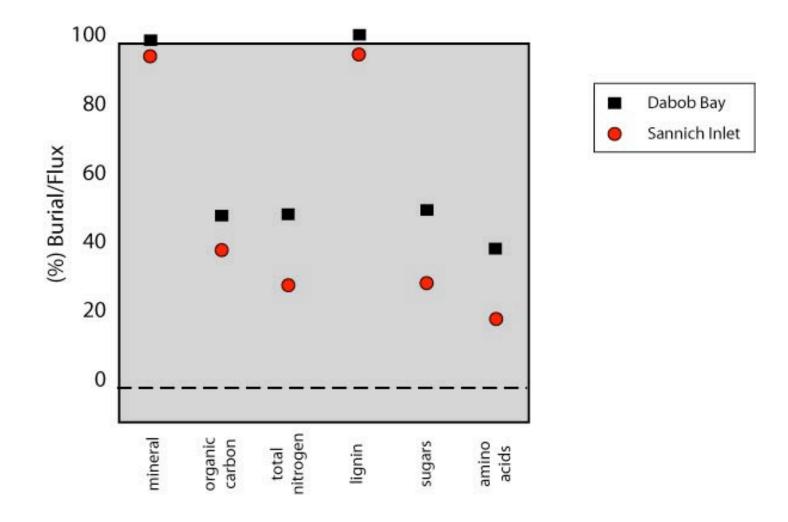


Kinetics of organic matter degradation and the multi "G" model

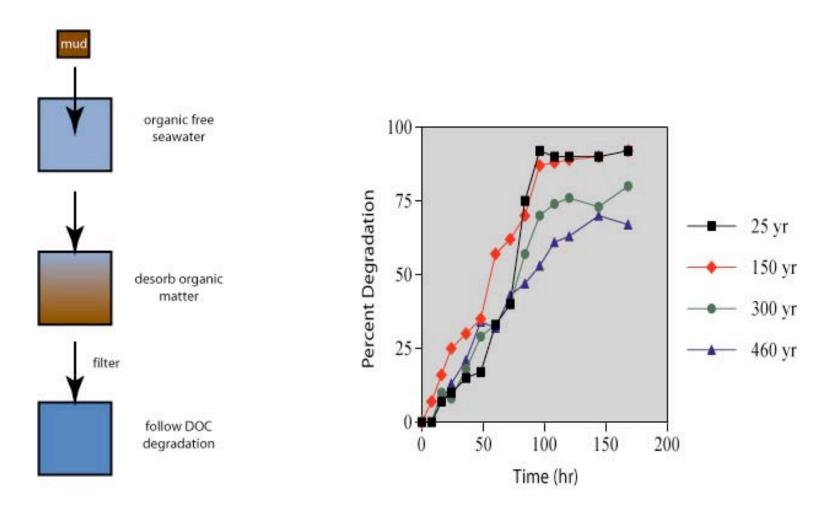




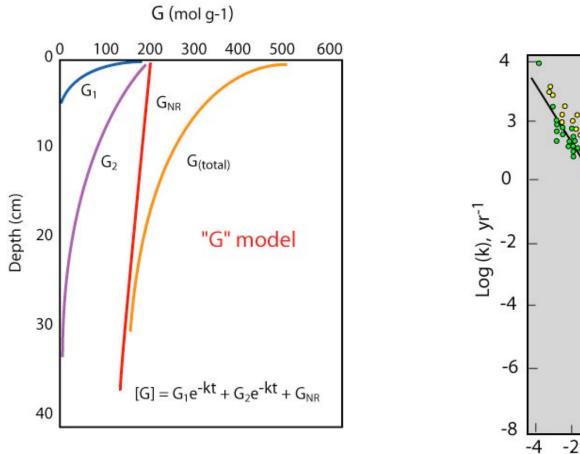
Selective preservation of organic matter in sediments

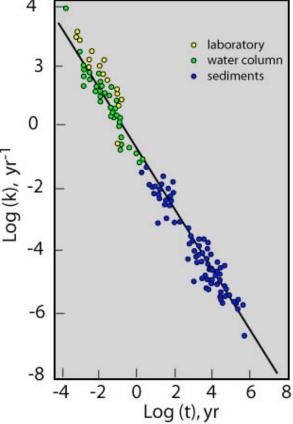


If selective preservation occurs, then "old", buried carbon should be recalcitrant. But is it ?



Is the "G" model just and observational artifact?

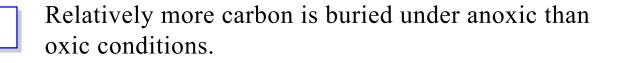


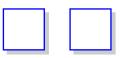






Productivity affects carbon burial efficiency.





Bigger molecules are degraded more slowly than smaller molecules.

The faster you bury carbon, the more likely it is to stay buried.

Figure by MIT OCW.

Carbon preservation summary

Ocean is >99.5% efficient at recycling C.

Annual production is about 50-70GT C yr⁻¹, of which 0.1-0.2% is buried.

Several factors affect C preservation.

organic matter production

oxygen

organic matter composition

It is difficult to isolate these factors from one another to elucidate underlying mechanisms