## Relationship between moist convection and Large-Scale flow

Boundary layer entropy equilibrium hypothesis (Raymond, 1995)



$$h\frac{\partial s_b}{\partial t} \approx 0 = F_s - \left(M_d + (1 - \sigma)w_d\right)\left(s_b - s_m\right)$$

Mass:

$$M_u - M_d - (1 - \sigma) w_d = w_b$$

$$\rightarrow M_u = w_b + \frac{F_s}{s_b - s_m}$$
(1)

Free troposphere heat balance:

$$\begin{pmatrix} M_u - M_d - w \end{pmatrix} \mathbf{S} = \dot{Q}_{cool},$$
$$\mathbf{S} \equiv c_p \frac{T}{\theta} \frac{\partial \theta}{\partial z}$$

Convective downdraft:

$$M_d = \left(1 - \varepsilon_p\right) M_u$$

$$\rightarrow \varepsilon_{p}M_{u} = w + \frac{\dot{Q}_{cool}}{\mathbf{S}}$$
(2)

Combine (1) and (2) Let  $w_b = \gamma w$ 

$$w = \frac{1}{1 - \gamma \varepsilon_p} \left[ \frac{\varepsilon_p F_s}{s_b - s_m} - \frac{\dot{Q}_{cool}}{\mathbf{S}} \right],$$
$$M_u = \frac{1}{1 - \gamma \varepsilon_p} \left[ \frac{F_s}{s_b - s_m} - \frac{\gamma \dot{Q}_{cool}}{\mathbf{S}} \right]$$

Note that  $M_u > w$ 

## Radiative-convective equilibrium: w=0

$$\rightarrow F_{s} = \frac{\dot{Q}_{cool} \left( s_{b} - s_{m} \right)}{\mathbf{S}\varepsilon_{p}},$$

$$M_{u} = \frac{\dot{Q}_{cool}}{\mathbf{S}\varepsilon_{p}}.$$

Define  

$$(F_s)_{eq} \equiv \frac{\dot{Q}_{cool}(s_b - s_m)}{\mathbf{S}\varepsilon_p}$$

Then  

$$w = \frac{\varepsilon_p}{1 - \gamma \varepsilon_p} \left[ \frac{F_s}{s_b - s_m} - \frac{(F_s)_{eq}}{(s_b - s_m)_{eq}} \right]$$
Surface fluxes:  $F_s \cong C_k |\mathbf{V}| (s_0^* - s_b)$ 

w > 0 if

• 
$$F_s > F_{eq}$$

- $(s_b s_m) < (s_b s_m)_{eq}$
- $Q_{cool} < (Q_{cool})_{eq}$

Note also that we must have  $M_{\mu} \ge 0$  so in circumstances under which (1) and (2) yield  $M_{\mu} < 0$ 

we take 
$$M_u = 0$$
,



radiative-subsidence balance

Weak Temperature Gradient Approximation (WTG) Sobel and Bretherton, 2000

- Ignore time dependence of T above PBL
- Determine w from aforementioned equations
- Determine vorticity from w
- Determine T by inverting balanced flow

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