Chapter 6

Isotropic homogeneous 3D turbulence

Turbulence was recognized as a distinct fluid behavior by Leonardo da Vinci more than 500 years ago. It is Leonardo who termed such motions "turbolenze", and hence the origin of our modern word for this type of fluid flow. But it wasn't until the beginning of last century that researchers were able to develop a rigorous mathematical treatment of turbulence. The first major step was taken by G. I. Taylor during the 1930s. Taylor introduced formal statistical methods involving correlations, Fourier transforms and power spectra into the turbulence literature. In a paper published in 1935 in the Proceedings of the Royal Society of London, he very explicitly presents the assumption that turbulence is a random phenomenon and then proceeds to introduce statistical tools for the analysis of homogeneous, isotropic turbulence. In 1941 the Russian statistician A. N. Kolmogorov published three papers (in Russian) that provide some of the most important and most-often quoted results of turbulence theory. These results, which will be discussed in some detail later, comprise what is now referred to as the K41 theory, and represent a major success of the statistical theories of turbulence. This theory provides a prediction for the energy spectrum of a 3D isotropic homogeneous turbulent flow. Kolmogorov proved that even though the velocity of an isotropic homogeneous turbulent flow fluctuates in an unpredictable fashion, the energy spectrum (how much kinetic energy is present on average at a particular scale) is predictable.

The spectral theory of Kolmogorov had a profound impact on the field and it still represents the foundation of many theories of turbulence. It it thus appropriate to start this course by introducing the concepts of 3D isotropic homogeneous turbulence and K41. It should however be kept in mind that 3D isotropic homogeneous turbulence is an idealization never encountered in nature. The challenge is then to understand what aspects of these theories apply to natural flows and what are pathological.



Courtesy of Andrew Ooi, University of Melbourne, Australia. Used with permission.

Figure 6.1: Isosurfaces of the the velocity gradient tensor used to visualize structures in computation of isotropic homogeneous 3D turbulence. The yellow surfaces represent flow regions with stable focus/stretching topology while the blue outlines of the isosurfaces show regions with unstable focus/contracting topology. 128^3 simulation with Taylor Reynolds number = 70.9. (Andrew Ooi, University of Melbourne, Australia, 2004, http://www.mame.mu.oz.au/fluids/).

A turbulent flow is said to be *isotropic* if,

- rotation and buoyancy are not important and can be neglected,
- there is no mean flow.

Rotation and buoyancy forces tend to suppress vertical motions, as we discuss later in the course, and create an anisotropy between the vertical and the horizontal directions. The presence of a mean flow with a particular orientation can also introduce anisotropies in the turbulent velocity and pressure fields.

A flow is said to be *homogeneous* if,

• there are no spatial gradients in any averaged quantity.

This is equivalent to assume that the statistics of the turbulent flow is not a function of space. An example of 3D isotropic homogeneous flow is shown in Fig. 6.1.

The theory of 3D isotropic homogeneous turbulence is based on the examination of the kinetic energy budget (potential energy is constant for flows with no buoyancy fluctuations),

$$\frac{\partial}{\partial t} \left(\frac{u_i^2}{2} \right) + \frac{\partial}{\partial x_j} \left(u_j \frac{u_i^2}{2} - \nu u_i \frac{\partial u_i}{\partial x_j} \right) = -\frac{\partial}{\partial x_i} \left(\frac{pu_i}{\rho_0} \right) - \nu \left(\frac{\partial u_i}{\partial x_j} \right)^2.$$
(6.1)

If we take an ensemble average of this equation under the assumptions of homogeneity we get,

$$\frac{\mathrm{dE}}{\mathrm{dt}} = -\epsilon,\tag{6.2}$$

where,

$$E \equiv \langle \frac{u_i^2}{2} \rangle, \qquad \epsilon \equiv \langle \nu \left(\frac{\partial u_i}{\partial x_j} \right)^2 \rangle.$$
 (6.3)

This equation state that the rate of change of turbulent kinetic energy E (TKE) is balanced by viscous dissipation ϵ . Such a balance cannot be sustained for long times - a source of kinetic energy is needed. However sources of TKE are typically not homogeneous: think of a stirrer or an oscillating boundary. We sidestep this contradiction by assuming that for large Reynolds numbers, although isotropy and homogeneity are violated by the mechanism producing the turbulence, they still hold at small scales and away from boundaries. Then the turbulence production can be represented simply by a forcing term F, assumed to be isotropic and homogeneous:

$$\frac{\mathrm{dE}}{\mathrm{dt}} = -\epsilon + F. \tag{6.4}$$

6.1 Kinetic Energy Spectra for 3D turbulence

Definition of KE in spectral space

For a flow which is homogeneous in space (i.e. statistical properties are independent of position), a spectral description is very appropriate, allowing us to examine properties as a function of wavelength. The total kinetic energy can be written replacing the ensemble average with a space average,

$$E = \frac{1}{2} \langle u_i^2 \rangle = \frac{1}{2} \frac{1}{V} \iiint u_i(\mathbf{x}) u_i(\mathbf{x}) \, \mathrm{d}\mathbf{x}, \tag{6.5}$$

where V is the volume domain. The spectrum $\phi_{i,j}(\mathbf{k})$ is then defined by,

$$E = \frac{1}{2} \iiint \phi_{i,i}(\mathbf{k}) \mathrm{d}\mathbf{k} = \iiint E(\mathbf{k}) \mathrm{d}\mathbf{k}$$
(6.6)

where $\phi_{i,j}(\mathbf{k})$ is the Fourier transform of the velocity correlation tensor $R_{i,j}(\mathbf{r})$,

$$\phi_{i,j}(\mathbf{k}) = \frac{1}{(2\pi)^3} \iiint R_{i,j}(\mathbf{r}) e^{-i\mathbf{k}\cdot\mathbf{r}} \mathrm{d}\mathbf{r}, \qquad R_{i,j}(\mathbf{r}) = \frac{1}{V} \iiint u_j(\mathbf{x}) u_i(\mathbf{x}+\mathbf{r}) \mathrm{d}\mathbf{x}. \quad (6.7)$$

 $R_{i,j}(\mathbf{r})$ tells us how velocities at points separated by a vector \mathbf{r} are related. If we know these two point velocity correlations, we can deduce $E(\mathbf{k})$. Hence the energy spectrum has the information content of the two-point correlation.

 $E(\mathbf{k})$ contains directional information. More usually, we want to know the energy at a particular scale $k = |\mathbf{k}|$ without any interest in separating it by direction. To find E(k), we integrate over the spherical shell of radius k (in 3-dimensions),

$$E = \iiint E(\mathbf{k}) \mathrm{d}\mathbf{k} = \int_0^\infty \left[\oint k^2 E(\mathbf{k}) \mathrm{d}\sigma \right] \mathrm{d}k = \int_0^\infty E(k) \mathrm{d}k, \tag{6.8}$$

where σ is the solid angle in wavenumber space, i.e. $d\sigma = \sin \theta_1 d\theta_1 d\theta_2$. We now define the isotropic spectrum as,

$$E(k) = \oint k^2 E(\mathbf{k}) \mathrm{d}\sigma = \frac{1}{2} \oint k^2 \phi_{i,i}(\mathbf{k}) \mathrm{d}\sigma.$$
(6.9)

For isotropic velocity fields the spectrum does not depend on directions, i.e. $\phi_{i,i}(\mathbf{k}) = \phi_{i,i}(k)$, and we have,

$$E(k) = 2\pi k^2 \phi_{i,i}(k).$$
(6.10)

Energy budget equation in spectral space

We have an equation for the evolution of the total kinetic energy E. Equally interesting is the evolution of E(k), the isotropic energy at a particular wavenumber k. This will include terms which describe the transfer of energy from one scale to another, via nonlinear interactions.

To obtain such an equation we must take the Fourier transform of the non-rotating, unstratified Boussinesq equations,

$$\frac{\partial u_i}{\partial t} - \nu \frac{\partial^2 u_i}{\partial x_j^2} = -u_j \frac{\partial u_i}{\partial x_j} - \frac{1}{\rho_0} \frac{\partial p}{\partial x_i}.$$
(6.11)

The two terms on the lhs are linear and are easily transformed into Fourier space,

$$\frac{\partial}{\partial t}u_i(\mathbf{x},t) \iff \frac{\partial}{\partial t}\hat{u}_i(\mathbf{k},t), \tag{6.12}$$

$$\nu \frac{\partial^2}{\partial x_j^2} u_i(\mathbf{x}, t) \iff -\nu k_j^2 \hat{u}_i(\mathbf{k}, t).$$
(6.13)

In order to convert the pressure gradient term, we first notice that taking the divergence of the Navier-Stokes equation we obtain,

$$\frac{\partial^2 p}{\partial x_i^2} = -\rho_0 \frac{\partial u_i}{\partial x_j} \frac{\partial u_j}{\partial x_i}.$$
(6.14)

Thus both terms on the rhs of eq. (6.11) involve the product of velocities. The convolution theorem states that the Fourier transform of a product of two functions is given by the convolution of their Fourier transforms,

$$\frac{1}{V} \iiint u_i(\mathbf{x}, t) u_j(\mathbf{x}, t) e^{i\mathbf{k}\cdot\mathbf{x}} d\mathbf{x} = \frac{1}{(2\pi)^3} \iiint \hat{u}_i(\mathbf{p}, t) \hat{u}_j(\mathbf{q}, t) \delta(\mathbf{p} + \mathbf{q} - \mathbf{k}) d\mathbf{p} d\mathbf{q}.$$
 (6.15)

Applying the convolution terms to the terms on the rhs we get, The two terms on the lhs are linear and are easily transformed in Fourier space,

$$u_j \frac{\partial u_i}{\partial x_j} \iff i \iiint q_j \hat{u}_j(\mathbf{p}, t) \hat{u}_i(\mathbf{q}, t) \delta(\mathbf{p} + \mathbf{q} - \mathbf{k}) \mathrm{d}\mathbf{p} \mathrm{d}\mathbf{q}, \tag{6.16}$$

$$p \iff \rho_0 \iiint \frac{p_i q_j}{k^2} \hat{u}_j(\mathbf{p}, t) \hat{u}_i(\mathbf{q}, t) \delta(\mathbf{p} + \mathbf{q} - \mathbf{k}) \mathrm{d}\mathbf{p} \mathrm{d}\mathbf{q}.$$
(6.17)

Plugging all these expressions in eq. (6.11) we obtain the Navier-Stokes equation in Fourier space,

$$\left(\frac{\partial}{\partial t} + \nu k^2\right)\hat{u}_i(\mathbf{k}, t) = -\mathrm{i} \iiint q_j \left(\delta_{i,m} - \frac{k_i p_m}{k^2}\right)\hat{u}_j(\mathbf{p}, t)\hat{u}_m(\mathbf{q}, t)\delta(\mathbf{p} + \mathbf{q} - \mathbf{k})\mathrm{d}\mathbf{p}\mathrm{d}\mathbf{q}$$
(6.18)

The term on the right hand side shows that the nonlinear terms involve triad interactions between wave vectors such that $\mathbf{k} = \mathbf{p} + \mathbf{q}$.

Now to obtain the energy equation we multiply e. (6.18) by $\hat{u}_i^*(\mathbf{k}, t)$ and we integrate over \mathbf{k} ,

$$\left(\frac{\partial}{\partial t} + 2\nu k^2\right)\phi_{i,i}(\mathbf{k},t) =$$

Re $\left[\iiint A_{ijm}(\mathbf{k},\mathbf{p},\mathbf{q})\hat{u}_i^*(\mathbf{k},t)\hat{u}_j(\mathbf{p},t)\hat{u}_m(\mathbf{q},t)\delta(\mathbf{p}+\mathbf{q}-\mathbf{k})\mathrm{d}\mathbf{p}\mathrm{d}\mathbf{q}\mathrm{d}\mathbf{k}\right].$ (6.19)

The terms on the rhs represent the triad interactions that exchange energy between $\hat{u}_i(\mathbf{k}, t)$, $\hat{u}_j(\mathbf{p}, t)$, and $\hat{u}_m(\mathbf{q}, t)$. The coefficient A_{ijm} are the coupling coefficient of each triad and depends only on the wavenumbers.

If pressure and advection were not present, the energy equation would reduce to,

$$\left(\frac{\partial}{\partial t} + 2\nu k^2\right)\phi_{i,i}(\mathbf{k},t) = 0, \qquad (6.20)$$

in which the wavenumbers are uncoupled. The solution to this equation is,

$$\phi_{i,i}(\mathbf{k},t) = \phi_{i,i}(\mathbf{k},0)e^{-\nu k^2 t}.$$
(6.21)

According to (6.21), the energy in wavenumber **k** decays exponentially, at a rate that increases with increasing wavenumber magnitude k. Thus viscosity damps the smallest spatial scale fastest.

We now restrict our analysis to isotropic velocity fields, so that we can use (6.10) and simplify (6.19),

$$\frac{\partial}{\partial t}E(k,t) = T(k,t) - 2\nu k^2 E(k,t), \qquad (6.22)$$

where T(k, t) comprises all triad interaction terms. If we examine the integral of this equation over all k,

$$\frac{\partial}{\partial t} \int_0^\infty E(k) \mathrm{d}k = \int_0^\infty T(k, t) \mathrm{d}k - 2\nu \int_0^\infty k^2 E(k) \mathrm{d}k, \tag{6.23}$$

and note that $-2\nu k^2 E(k)$ is the Fourier transform of the dissipation term, then we see that the equation for the total energy budget in (6.2), is recovered only if,

$$\int_0^\infty T(k,t)\mathrm{d}k = 0. \tag{6.24}$$

Hence the nonlinear interactions transfer energy between different wave numbers, but do not change the total energy.

Now, adding a forcing term to the energy equation in k-space we have the following equation for energy at a particular wavenumber k,

$$\frac{\partial}{\partial t}E(k,t) = T(k,t) + F(k,t) - 2\nu k^2 E(k,t), \qquad (6.25)$$

where F(k,t) is the forcing term, and T(k,t) is the **kinetic energy transfer**, due to nonlinear interactions. The **kinetic energy flux** through wave number k is $\Pi(k,t)$, defined as,

$$\Pi(k,t) = \int_{k}^{\infty} T(k',t) \mathrm{d}k', \qquad (6.26)$$

or

$$T(k,t) = -\frac{\partial \Pi(k,t)}{\partial k}.$$
(6.27)

For stationary turbulence,

$$2\nu k^2 E(k) = T(k) + F(k).$$
(6.28)

Remembering that the total dissipation rate is given by,

$$\epsilon = \int_0^\infty 2\nu k^2 E(k) \mathrm{d}k \tag{6.29}$$

and that the integral of the triad interactions over the whole k-space vanishes, we have,

$$\epsilon = \int_0^\infty F(k) \mathrm{d}k. \tag{6.30}$$

The rate of dissipation of energy is equal to the rate of injection of energy.

If the forcing F(k) is concentrated on a narrow spectral band centered around a wave number k_i , then for $k \neq k_i$,

$$2\nu k^2 E(k) = T(k). (6.31)$$

In the limit of $\nu \to 0$, the energy dissipation becomes negligible at large scales. Thus there must be an intermediate range of scales between the forcing scale and the scale where viscous dissipation becomes important, where,

$$2\nu k^2 E(k) = T(k) \approx 0.$$
 (6.32)

Notice that ϵ must remain nonzero, for nonzero F(k), in order to balance the energy injection. This is achieved by $\int_0^\infty k^2 E(k) dk \to \infty$, i.e. the velocity fluctuations at small scales increase.

Then we find the energy flux in the limit $\nu \to 0$,

$$\Pi(k) = 0, : k < k_i$$

$$\Pi(k) = \epsilon : k > k_i$$
(6.33)

Hence at vanishing viscosity, the kinetic energy flux is constant and equal to the injection rate, for wavenumbers greater than the injection wavenumber k_i . The scenario is as follows. (a) Energy is input at a rate ϵ at a wavenumber k_i . (b) Energy is fluxed to higher wavenumbers at a rate ϵ trough triad interactions. (c) Energy is eventually dissipated at very high wavenumbers at a rate ϵ , even in the limit of $\nu \to 0$.

The statement that triad interactions produce a finite energy flux ϵ toward small scales does not mean that all triad interactions transfer energy exclusively toward small scales. Triad interactions transfer large amounts of energy toward both large and small scales. On average, however, there is an excess of energy transfer toward small scales given by ϵ .

Kolmogorov spectrum

Kolmogorov's 1941 theory for the energy spectrum makes use of the result that ϵ , the energy injection rate, and dissipation rate also controls the flux of energy. Energy flux is independent of wavenumber k, and equal to ϵ for $k > k_i$. Kolmogorov's theory assumes the injection wavenumber is much less than the dissipation wavenumber $(k_i \ll k_d, \text{ or large Re})$. In the intermediate range of scales $k_i \ll k_d$ neither the

forcing nor the viscosity are explicitly important, but instead the energy flux ϵ and the local wavenumber k are the only controlling parameters. Then we can express the energy density as

$$E(k) = f(\epsilon, k) \tag{6.34}$$

Now using dimensional analysis:

Quantity	Dimension
Wavenumber k	1/L
Energy per unit mass E	$U^2 \sim L^2/T^2$
Energy spectrum $E(k)$	$EL \sim L^3/T^2$
Energy flux ϵ	$E/T \sim L^2/T^3$

In eq. (6.34) the lhs has dimensionality L^3/T^2 ; the dimension T^{-2} can only be balanced by $\epsilon^{2/3}$ because k has no time dependence. Thus,

$$E(k) = \epsilon^{2/3} g(k). \tag{6.35}$$

Now g(k) must have dimensions $L^{5/3}$ and the functional dependence we must have, if the assumptions hold, is,

$$E(k) = C_K \epsilon^{2/3} k^{-5/3} \tag{6.36}$$

This is the famous Kolmogorov spectrum, one of the cornerstone of turbulence theory. C_K is a universal constant, the Kolmogorov constant, experimentally found to be approximately 1.5. The region of parameter space in k where the energy spectrum follows this $k^{-5/3}$ form is known as the **inertial range**. In this range, energy **cascades** from the larger scales where it was injected ultimately to the dissipation scale. The theory assumes that the spectrum at any particular k depends only on spectrally local quantities - i.e. has no dependence on k_i for example. Hence the possibility for long-range interactions is ignored.

We can also derive the Kolmogorov spectrum in a perhaps more physical way (after Obukhov). Define an eddy turnover time $\tau(k)$ at wavenumber k as the time taken for a parcel with energy E(k) to move a distance 1/k. If $\tau(k)$ depends only on E(k) and k then, from dimensional analysis,

$$\tau(k) \sim \left[k^3 E(k)\right]^{-1/2}$$
 (6.37)

The energy flux can be defined as the available energy divided by the characteristic time τ . The available energy at a wavenumber k is of the order of kE(k). Then we have,

$$\epsilon \sim \frac{kE(k)}{\tau(k)} \sim k^{5/2} E(k)^{3/2},$$
(6.38)

and hence,

$$E(k) \sim \epsilon^{2/3} k^{-5/3}.$$
 (6.39)

Characteristic scales of turbulence

Kolmogorov scale

We have shown that viscous dissipation acts most efficiently at small scales. Thus above a certain wavenumber k_d , viscosity will become important, and E(k) will decay more rapidly than in the inertial range. The regime $k > k_d$ is known as the *dissipation* range. A simple scaling argument for k_d can be made by assuming that the spectrum follows the inertial scaling until k_d and then drops suddenly to zero because of viscous dissipation. In reality the transition between the two regimes is more gradual, but this simple model predicts k_d quite accurately. First we assume,

$$E(k) = C_K \epsilon^{2/3} k^{-5/3}, \qquad k_i < k < k_d, E(k) = 0, \qquad k > k_d.$$
(6.40)

Substituting (6.29), and integrating between k_i and k_d we find,

$$k_d \sim \left(\frac{\epsilon^{1/4}}{\nu^{3/4}}\right). \tag{6.41}$$

The inverse $l_d = 1/k_d$ is known as the *Kolmogorov scale*, the scale at which dissipation becomes important.

$$l_d \sim \left(\frac{\nu^{3/4}}{\epsilon^{1/4}}\right) \tag{6.42}$$

Integral scale

At the small wavenumber end of the spectrum, the important lengthscale is l_i , the integral scale, the scale of the energy-containing eddies. $l_i = 1/k_i$. We can evaluate l_i in terms of ϵ . Let us write,

$$U^2 = 2 \int_0^\infty E(k) \mathrm{d}k \tag{6.43}$$

and substituting for E(k) from (6.36),

$$U^{2} \sim 2 \int_{0}^{\infty} C_{K} \epsilon^{2/3} k^{-5/3} \mathrm{d}k \sim 3C_{K} \epsilon^{2/3} k_{i}^{-2/3}.$$
 (6.44)

Then,

$$k_i \sim \frac{\epsilon}{U^3} \tag{6.45}$$

so that $l_i \sim U^3/\epsilon$. Then the ratio of maximum and minimum dynamically active scales,

$$\frac{l_i}{l_d} = \frac{k_d}{k_i} \sim \frac{U^3}{\epsilon^{3/4} \nu^{3/4}} \sim \left(\frac{Ul_i}{\nu}\right)^{3/4} \sim Re_{l_i}^{3/4}.$$
(6.46)

where Re_{l_i} is the *integral Reynolds number*. Hence in K41 the inertial range spans a range of scales growing as the (3/4)th power of the integral Reynolds number. It follows that if we want to describe such a flow accurately in a numerical simulation on a uniform grid, the minimum number of points per integral scale is $N \sim Re_{l_i}^{9/4}$. One consequence is that the storage requirements of numerical simulations scale as $Re_{l_i}^{9/4}$. Since the time step has usually to be taken proportional to the spatial mesh, the total computational work needed to integrate the equations for a fixed number of large eddy turnover times grows as $Re_{l_i}^3$. This shows that progress in achieving high Re_{l_i} simulations is very slow.

6.1.1 Kolmogorov in physical space

Kolmogorov formulated his theory in physical space, making predictions for S_p , the longitudinal velocity structure function of order p,

$$\delta v_r = \left[\mathbf{u}(\mathbf{x} + \mathbf{r}, t) - \mathbf{u}(\mathbf{x}, t) \right] \cdot \frac{\mathbf{r}}{r}$$
(6.47)

$$S_p = \langle |\delta v_r|^p \rangle. \tag{6.48}$$

For homogeneous isotropic turbulence the structure function depends only on the magnitude of \mathbf{r} , i.e. $S_p = S_p(r)$. Under the assumptions described above, i.e. that at a scale r, S_p depends only on the energy flux ϵ , and the scale r, dimensional analysis can be sued to predict that,

$$S_p(r,t) = C_p(\epsilon r)^{p/3} \tag{6.49}$$

where C_p is a constant. In particular $S_2 \sim (\epsilon r)^{2/3}$. The second order structure function is related to the energy spectrum for an isotropic homogeneous field,

$$S_{2} = \langle (\delta v_{r})^{2} \rangle = \langle (u_{//}(\mathbf{x} + \mathbf{r}, t) - u_{//}(\mathbf{x}, t))^{2} \rangle$$

$$= 2 \langle u_{//}(\mathbf{x} + \mathbf{r}, t) u_{//}(\mathbf{x}, t) \rangle + 2 \langle |u_{//}(\mathbf{x}, t)|^{2} \rangle$$

$$= 2 \int (1 - e^{i\mathbf{k} \cdot \mathbf{r}}) \phi(\mathbf{k}, t) d\mathbf{k}$$

$$= 4 \int_{0}^{\infty} E(k, t) \left(1 - \frac{\sin(kr)}{kr} \right) dk. \qquad (6.50)$$

If we substitute for E(k) from the Kolmogorov spectrum, and assume this applies from $k \gg r^{-1}$ then,

$$\langle (\delta v_r)^2 \rangle \sim C_K (\epsilon r)^{2/3}.$$
 (6.51)

Hence the Kolmogorov $k^{-5/3}$ spectrum is consistent with the second order structure function of the form $r^{2/3}$. (Note that S_2 is only finite if E(k,t) has the form k^{-n} where 1 < n < 3.)

6.1.2 3D homogeneous isotropic turbulence in geophysical flows

The assumptions of homogeneity, stationarity and isotropy as employed by Taylor and Kolmogorov have permitted tremendous advances in our understanding of turbulence. In particular the theory of Kolmogorov remains an outstanding example of what we mean by emergence of statistical predictability in a chaotic system. However this a a course on geophysical turbulence and we must ultimately confront the fact that real world physical flows rarely conform to our simplifying assumptions. In geophysical turbulence, statistical symmetries are upset by a complex interplay of effects. Here we focus on three important class of phenomena that modify small-scale turbulence in the atmosphere and ocean: large-scale shear, stratification, and boundary proximity. In a few weeks we will consider the role of rotation.

A very nice description of the effect of shear, stratification, and boundary proximity on small-scale geophysical turbulence is given in the review article on "3D Turbulence" by Smyth and Moum included as part of the reading material. The student should read that paper before proceeding with this chapter. Here we only provide some additional comments on the definition of the Ozimdov scale.

Ozmidov scale

In geophysical flows 3D turbulence can be a reasonable approximation at scales small enough that buoyancy and rotation effects can be neglected. Stratification becomes important at scales smaller than rotation and it is therefore more important in setting the upper scales at which 3D arguments hold. Stratification affects turbulence when the Froude number Fr = U/(NH) < 1, where U is a typical velocity scale, and H a typical vertical length scale of the motion. For large Fr, the kinetic energy of the motion is much larger than the potential energy changes involved in making vertical excursions of order H. For small Fr, the stratification suppresses the vertical motion because a substantial fraction of kinetic energy must be converted to potential energy when a parcel moves in the vertical.

We can define a characteristic scale l_B at which overturning is suppressed by the buoyancy stratification as follows. The velocity associated with a particular length scale l in high Reynolds number isotropic 3D turbulence scales like,

$$u^2 \sim \epsilon^{2/3} k^{-2/3} \qquad \Longleftrightarrow \qquad u \sim (l\epsilon)^{1/3}.$$
 (6.52)

Vertical motion at length scale l will be suppressed by the stratification when the local Froude number $Fr_l = 1$. If we define the length scale at which this suppression occurs as l_B then,

$$\frac{u_B}{Nl_B} = \frac{(l_B\epsilon)^{1/3}}{Nl_B} = 1 \qquad \Longrightarrow \qquad l_B = \left(\frac{\epsilon}{N^3}\right)^{1/2} \tag{6.53}$$

where l_B is known as the Ozmidov scale.

In stratified geophysical flows, we have a scenario in which a regime transition occurs at l_B .

- $l < l_B$: Fully 3D, isotropic turbulence. In this regime stratification can be neglected, and an inertial range may exist, if $l_d << l_B$, i.e. $\epsilon/(\nu N^2) >> 1$.
- $l > l_B$: Stratification influenced regime. In this regime ϵ is no longer constant with wave number, since some kinetic energy is lost through conversion to potential energy. 3D turbulence is replaced by motion controlled by the buoyancy stratification: either internal waves, or a quasi-2-dimensional turbulence, often described as "pancake turbulence", characterized by strong vortical motions in decoupled horizontal layers.

(For more on stratified turbulence, see Lesieur Ch XIII, Metais and Herring, 1989: Numerical simulations of freely evolving turbulence in stably stratified fluids. *J. Fluid Mech.*, **239**. Fincham, Maxworthy and Spedding, 1996: Energy dissipation and vortex structure in freely-decaying stratified grid turbulence. *Dyn. Atmos. Oceans*, **23**, 155-169.)

Further reading: Lesieur, Ch V, VI; Tennekes and Lumley, Ch 8; Frisch, Ch 5, 6, 7, 8.